
Coupling the Dynamics of Boundary Layers and Evolutionary Dunes



Great Sand Dunes Nat. Park, CO,USA (P.S. & P.O.)

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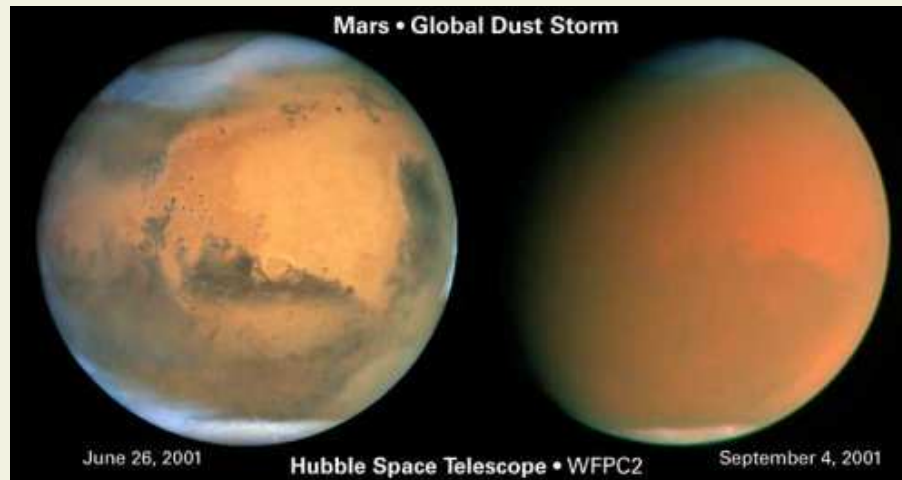
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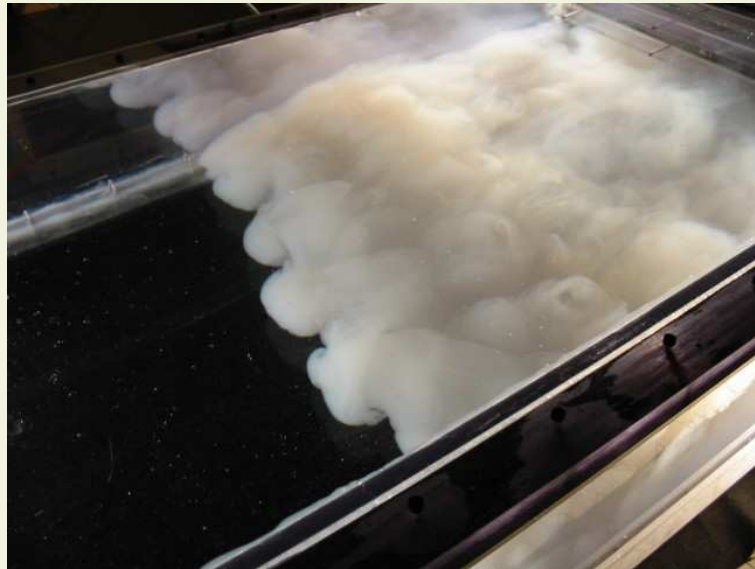
1. Sediment transport and landforms. *(a) Keys*

- Scales of the problem
 - Planetary scale



Mars before and after the great Martian dust storm of 2001
(From MGS, Mars Global Surveyor, and Hubble Space Telescope)

- Local scale
 - * Sediment dominated currents



Turbidity Currents. Lab tank
(From www.physics.utoronto.ca)

* Single morphologies: Barchan dunes



Simple barchan (1), Large simple barchan (2), Megabarchan (3) and ripple patterns (4). Location: 8 km. SE Chimbote, Perú.
(by J. McCauley, USGS, 1971)

* Complex morphologies: Dune fields

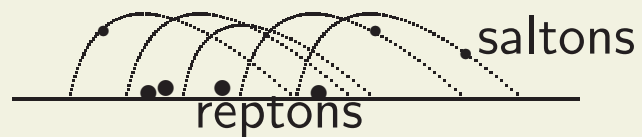


Great Sand Dunes Nat. Park, CO, USA
(by P. Smolarkiewicz & P. Ortiz, 2003)

- Micro-scale dynamics: Saltation, reptation and suspension



wind direction



- **Turbulence**
- **Separating SPBL**
- **Active role of landform:** Intricate geometry time - dependent boundary forcing. Active in transport: slopes and avalanches

1. Sediment transport and landforms. *(b) Solutions*

- QSA: Quasi steady approximations
 - Extreme wind scenarios: Full coupling

- Pre requisites of the model:
 - Time dependent curvilinear coordinate transformation
 - LES, Smagorinsky type SGS model.
 - Sediment transport model:
 - * Accomodated as a Convection Diffusion PDE
 - * Saltation fluxes as convective fluxes
 - * Sand avalanches as diffusive fluxes.(Anisotropic, inhomogeneous)

2.A Fluid Model

- Formulation in generalized time dependent coordinates

$$(\bar{t}, \bar{\mathbf{x}}) \equiv (t, \mathcal{F}(t, \mathbf{x})) ,$$

- Incompressible Boussinesq eqs, neutrally stratified flows, phys. space,

$$\nabla \cdot (\rho_o \mathbf{v}) = 0 ,$$

$$d\mathbf{v}/dt = -\nabla\pi' + \mathcal{D}(\mathcal{E}, \nabla\mathbf{v}) ,$$

$$d\mathcal{E}/dt = \mathcal{S}(\mathcal{E}) ,$$

-
- Dependence of \bar{x} , \bar{y} , \bar{z} , \bar{t} on (x, y, z, t) :

$$\bar{t} = t, \quad \bar{x} = x, \quad \bar{y} = y$$

$$\bar{z} := H_0 \frac{z - h(x, y, t)}{H_0(x, y, t) - h(x, y, t)},$$

h : lower surface elevations, physical domain,

H_0 : vertical extend of the transformed model domain.

- Solution of the sediment motion model: h : solid/fluid interface profile.

2.B Sediment Motion

- Evolution of the interface:

$$\rho_s \frac{\partial h}{\partial t} + \nabla_H \cdot \mathbf{q} = 0 ,$$

$\rho_s = \rho_m(1 - \lambda)$: bulk density of the sediment. ρ_m , λ : density of the grain material, porosity (volume of voids/total volume); ∇_H : (horizontal, physical space); \mathbf{q} : vertically integrated sediment mass flux.

- Saltation transport:

$$\mathbf{q}_S = C \frac{\rho}{g} \mathbf{u}_* \|\mathbf{u}_*\|^2 \max \left(0, 1 - \frac{u_\tau}{\|\mathbf{u}_*\|} \right) ,$$

C empirical coefficient depending upon the normalized grain size; ρ : density of the air; $g = |\mathbf{g}|$; $\mathbf{u}_* \equiv u_* \mathbf{v} / \|\mathbf{v}\|^{-1}$, friction velocity $u_* = \sqrt{\rho^{-1} \tau_w}$; τ_w : wall shear stress;

- u_τ : threshold value of u_* . Dependence on the slope:

$$u_\tau = \sqrt{\frac{\sin \theta}{\tan \alpha} \cos \gamma + \sqrt{\frac{\sin^2 \theta}{\tan^2 \alpha} (\cos^2 \gamma - 1) + \cos^2 \theta}} u_{\tau 0} ,$$

θ : local slope angle; α : angle of friction; γ : angle between local wind and slope. $u_{\tau 0}$: horizontal bed.

- Avalanche transport



Local sand avalanches beneath the brink of a dune
Great Sand Dunes Nat. Park, CO, USA
(by P. Smolarkiewicz & P. Ortiz, 2003)

Diffusion fluxes, anisotropic inhomogeneous diffusion coefficient $\mathcal{K}(\mathbf{x}, t)|_{\bar{z}=0}$ depending critically on the local slope

$$\mathbf{q}_A = -\rho_s \mathcal{K} \nabla_H h .$$

$$\mathcal{K} := \frac{\Lambda^2}{\Upsilon} \frac{1 + \operatorname{sgn}(\|\nabla_H h\| - s_C)}{2} ,$$

Λ and Υ : characteristic length and time scales.

- Total flux as advection-diffusion equation

$$\frac{\partial h}{\partial t} + \nabla_H \bullet \mathbf{U} h = \nabla_H \bullet \mathcal{K} \nabla_H h; \quad \mathbf{U} := \frac{\mathbf{q}_S}{\rho_s h} .$$

\mathbf{U} : average over potentially mobilized sand layer.

3. Numerical model

- Eulerian conservation law

$$\frac{\partial \rho^* \psi}{\partial \bar{t}} + \bar{\nabla} \bullet (\bar{\mathbf{V}}^* \psi) = \rho^* R ,$$

ψ : components of \mathbf{v} or \mathcal{E} , $\bar{\mathbf{V}}^* \equiv \rho^* \bar{\mathbf{v}}^*$, R : rhs, $\rho^* \equiv \rho_o \bar{G}$, \bar{G} : Jacobian of the transformation.

- NFT algorithm (second-order accuracy)

$$\psi_{\mathbf{i}}^{n+1} = \frac{\rho^{*n}}{\rho^{*n+1}} \mathcal{A}_{\mathbf{i}}(\tilde{\psi}, \bar{\mathbf{V}}^{*n+1/2}, \delta t) + 0.5 \delta t R_{\mathbf{i}}^{n+1} ;$$

ψ_i^{n+1} : solution at $(\bar{t}^{n+1}, \bar{\mathbf{x}}_i)$. $\tilde{\psi} \equiv \psi^n + 0.5\delta t R^n$.

For \mathcal{A} : fully second-order-accurate multidim. MPDATA advection scheme (PS & LM 98, ...).

- SGS forcings in R explicit.
- Solution for velocity and pressure (compact form):

$$\mathbf{v}_i = \hat{\mathbf{v}}_i - 0.5\delta t \left(\tilde{\mathbf{G}} \bar{\nabla} \pi' \right)_i ,$$

$$\left[\frac{\delta t}{\rho^*} \bar{\nabla} \bullet \rho^* \tilde{\mathbf{G}}^T \left(\hat{\mathbf{v}} - \tilde{\mathbf{G}} \bar{\nabla} \pi'' \right) \right]_i = 0 ,$$

$$\tilde{\mathbf{G}}^T \left(\hat{\mathbf{v}} - \tilde{\mathbf{G}} \bar{\nabla} \pi'' \right) \equiv \bar{\mathbf{v}}^s \text{ (J.P. \& P.S., 2003).}$$

Boundary value problem: Preconditioned GCR(k) algorithm.

-
- Updated pressure, updated solenoidal velocity: updated physical and contravariant velocity components using transformations.
 - Sediment Transport numerical model:
 - Explicit integration to $\mathcal{O}(\delta t^2)$ using the NFT

$$h_{\mathbf{j}}^{n+2} = \frac{\overline{G_{xy}}^n}{\overline{G_{xy}}^{n+2}} \mathcal{A}_{\mathbf{j}}^H(\tilde{h}, \overline{\mathbf{U}}^{*n}, 2\delta t) ;$$

$\mathcal{A}_{\mathbf{j}}^H$: nonoscillatory horizontal-advection operator.

$\tilde{h} \equiv h^n + 2\delta t \mathcal{L}(\mathcal{K}^n, h^n)$, \mathcal{L} : Laplacian

for all $\mathbf{j} = (i, j)$.

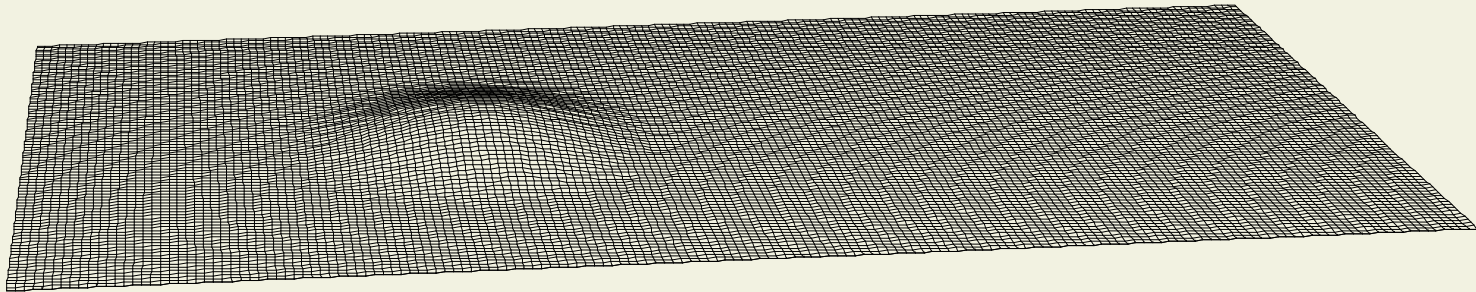
- Preventing negative h by limiting advective fluxes.

4. Results

- Neutrally stratified, nonrotating Boussinesq atmosphere. Uniform ambient wind $\mathbf{v}_e = (11, 0, 0) \text{ ms}^{-1}$
- Cartesian model domain $L_x \times L_y \times L_z = 46h_o \times 32h_o \times 5.3h_o$. (for $h_o = 7.5 \text{ m}$, $340 \times 180 \times 40 \text{ m}^3$).
- Lower boundary at $t = 0$: cosine sandpile of height h_o (range (0.5 – 7.5) m. Half-width a , $h_o/a \approx 0.15$, centered at $(x_o, y_o) = (L_x/3, L_y/2)$, a range: $\approx(3 - 50) \text{ m}$.

$$h(\mathbf{x}, t = 0) = \begin{cases} h_o \cos^2\left(\frac{\pi r}{2a}\right) + h_b & \text{if } r/a \leq 1, \\ h_b & \text{if } r/a > 1. \end{cases}$$

Initial conditions. Profile



$r \equiv \sqrt{(x - x_o)^2 + (y - y_o)^2}$. h_b : thickness of the sand layer: 0 or 5 m.

- Time scales: PBL flows $\mathcal{O}(10^3)$ s. Dune evolution $\mathcal{O}(10^6)$ s: Rescaling C : 1440 (minutes per day).
- Upper boundary: rigid lid. Lateral: Open and periodic (streamwise and spanwise).
- Initial condition: potential flow.
- Regular mesh (in the transformed space) $171 \times 91 \times 41$ and $333 \times 119 \times 41$.

-
- Surface drag coefficient in the Smagorinsky-type turbulence subgrid-scale model: $C_D = 0.01$.
 - Spatial/temporal scales \mathbf{q}_A : $\Lambda = 0.25 \min(\delta x, \delta y)$, $\Upsilon = \delta t$.

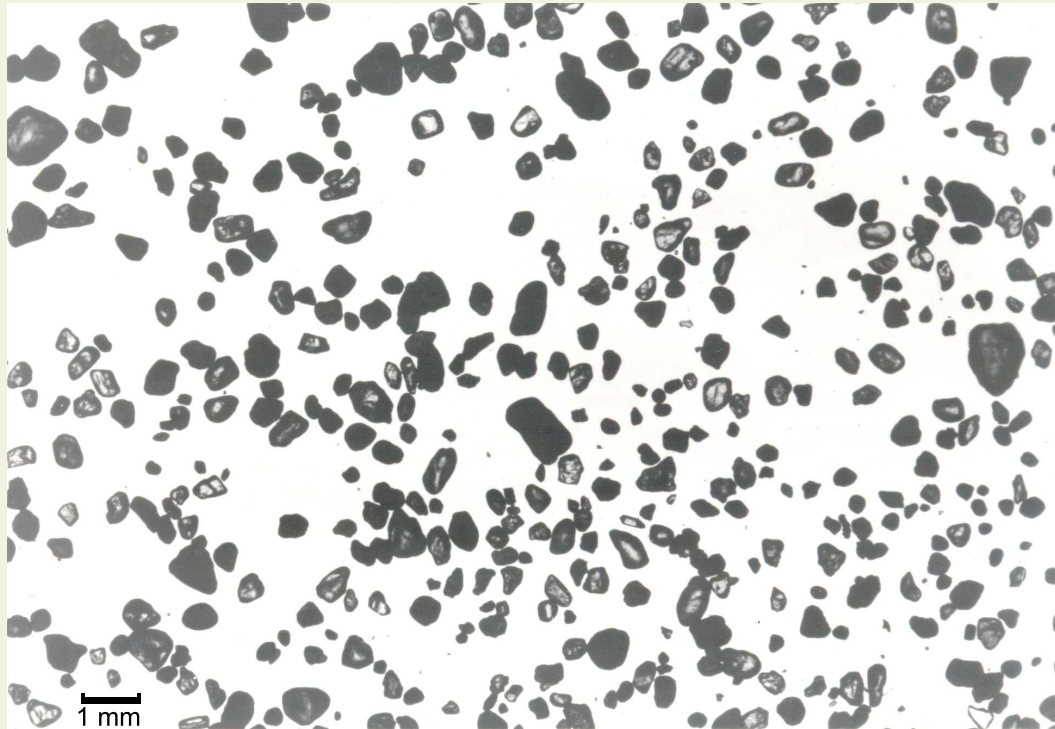
- Friction velocity:

$$\mathbf{u}_* = \kappa \frac{(\mathbf{v} - \mathbf{v} \cdot \mathbf{n})|_{z_\Delta}}{\ln(z_\Delta/z_0)}$$

z_Δ : surface-adjacent level. z_0 : equivalent roughness length (flow-to-grains momentum transfer) $z_0 = 0.001$ m. $\kappa = 0.41$ von Karman constant.

- Sediment transport parameters: $\rho_m = 2650$ kg/m³ (quartz). $\lambda = 0.5$. $u_{\tau 0} = 0.22$ ms⁻¹.

- Collected sand sample. Sand Dunes Nat. Park



h_o [m]	H [m]	Δt [s]	t_f/T_o
7.5	9.5	0.05	232848
6.0	7.6	0.03	174636
3.0	3.7	0.02	232848
1.5	1.6	0.01	209088
1.0	1.0	0.005	121176
0.7 †	0.6	0.002	105494
0.5 †	0.4	0.002	105494

Experiments. h_o : initial heights. H : final max. height. Δt : time step.

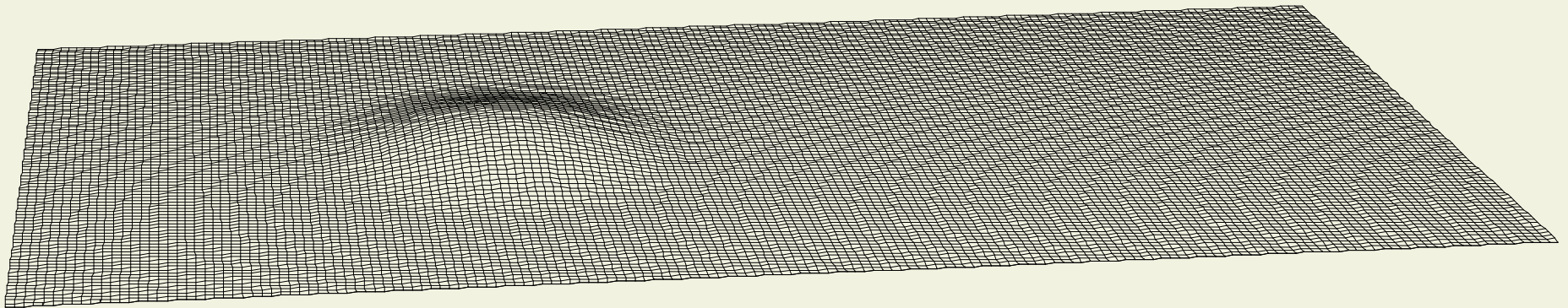
$h_o/a \approx 0.15$; $h_b = 0$; $L_x/h_o \approx 46$; $L_y/h_o \approx 32$; $L_z/h_o \approx 5.3$;

$171 \times 119 \times 41$ grid points; $\Delta x \approx 0.27h_o$, $\Delta y \approx 0.27h_o$, $\Delta z \approx 0.132h_o$;

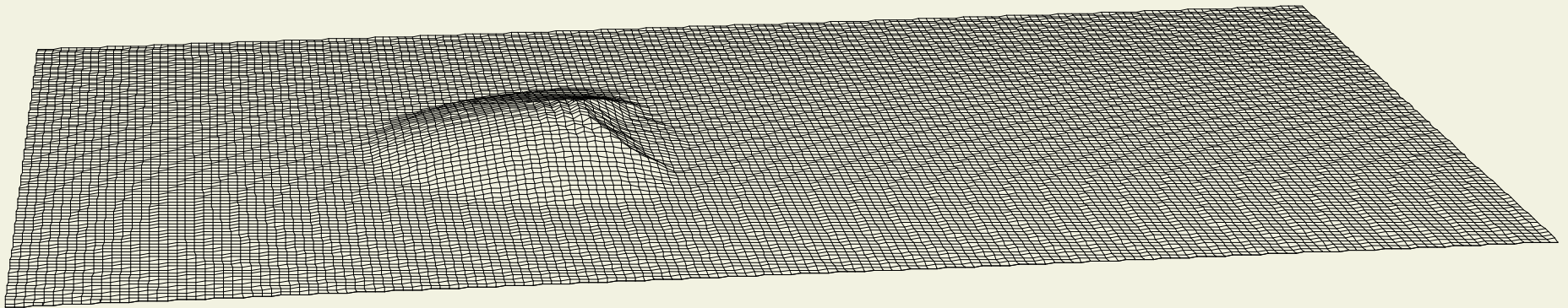
†: $h_o/a \approx 0.15$; $h_b = 0$; $L_x/h_o \approx 92$; $L_y/h_o \approx 32$; $L_z/h_o \approx 5.3$;

$333 \times 119 \times 41$ grid points; $\Delta x \approx 0.27h_o$, $\Delta y \approx 0.27h_o$, $\Delta z \approx 0.132h_o$.

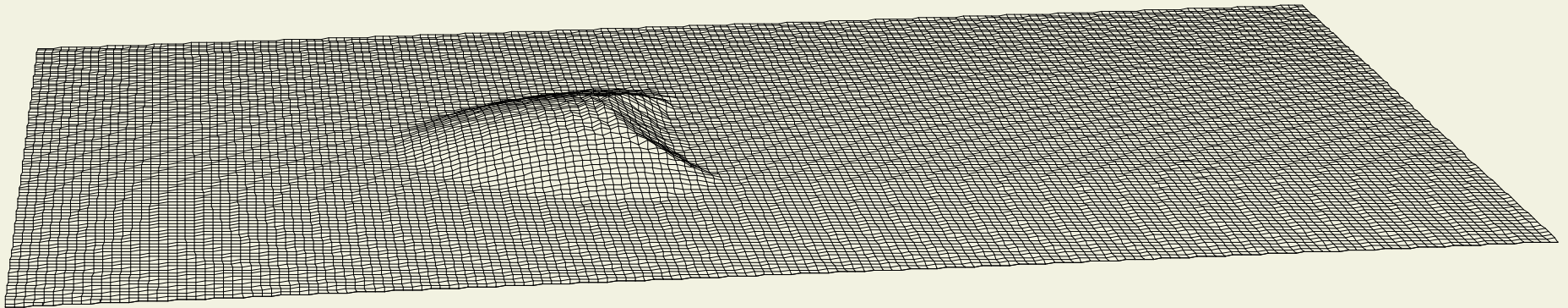
Non erodible substrate profile



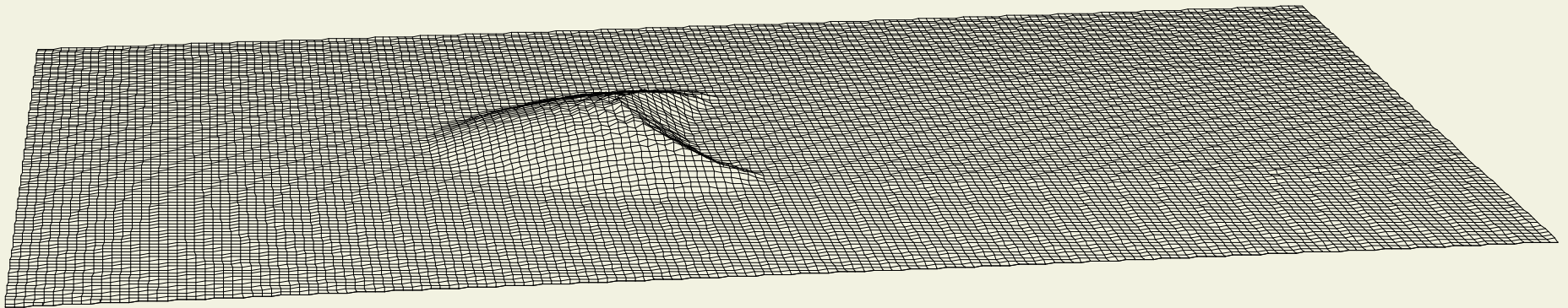
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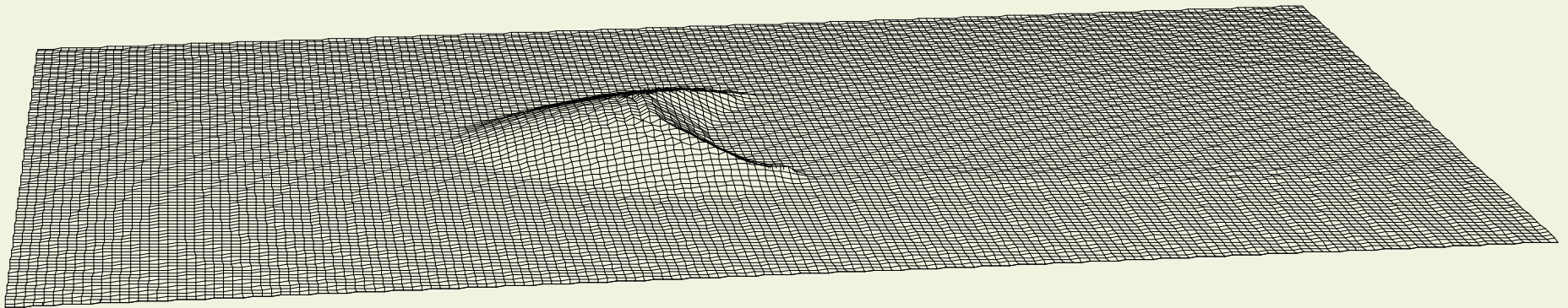
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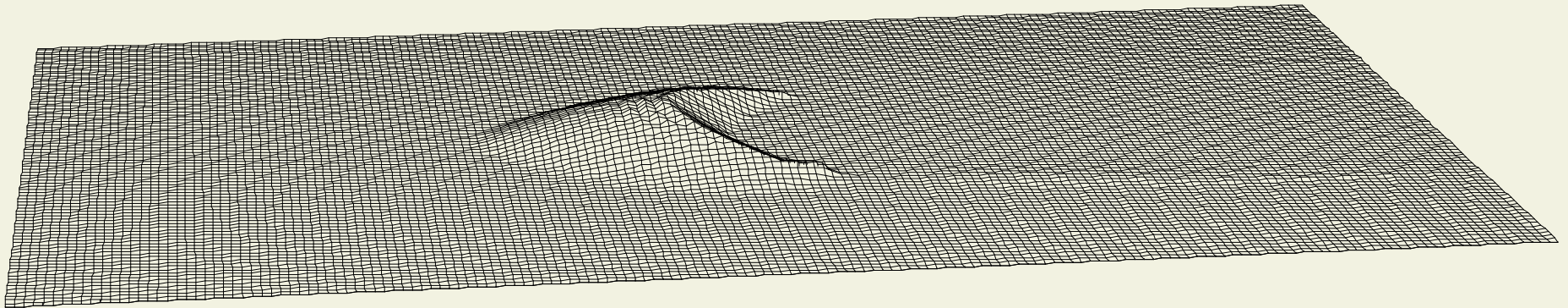
Non erodible substrate profile



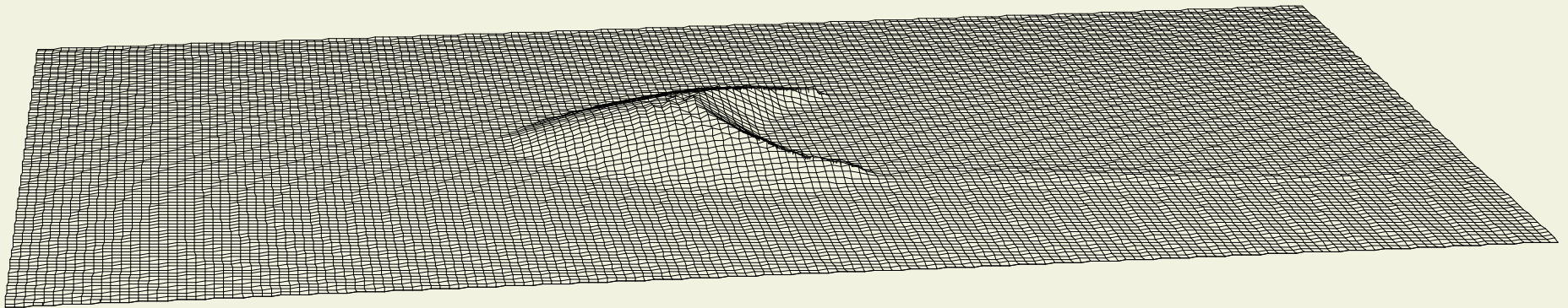
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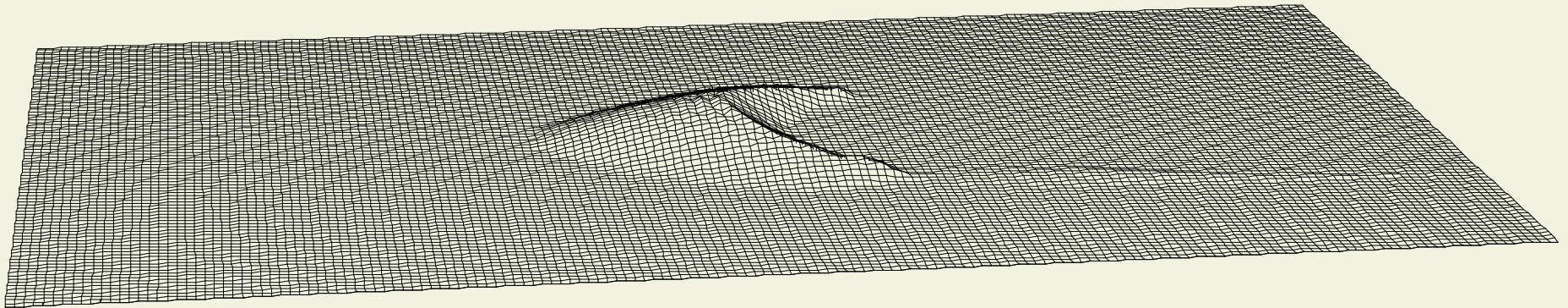
Non erodible substrate profile



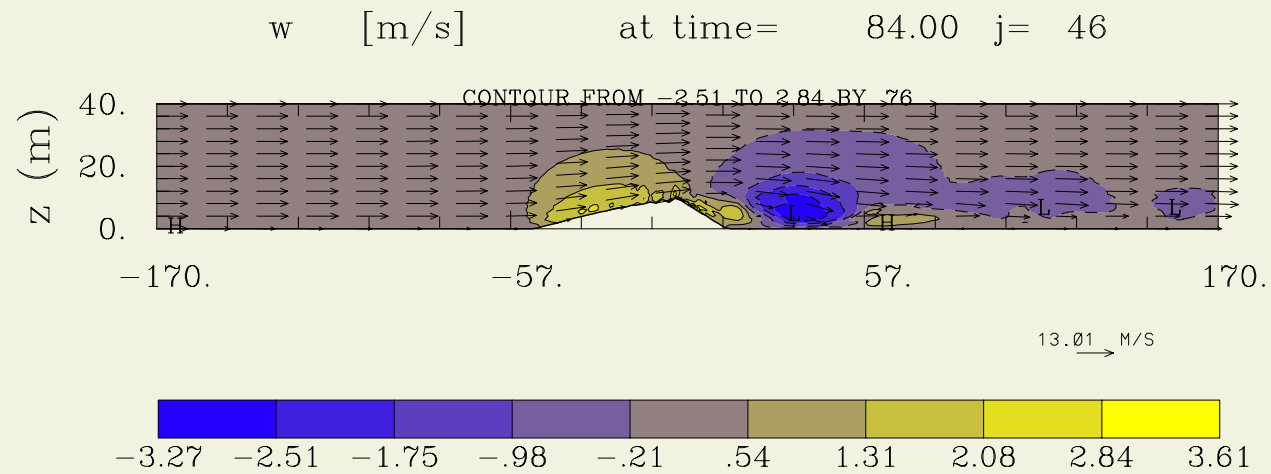
Non erodible substrate profile



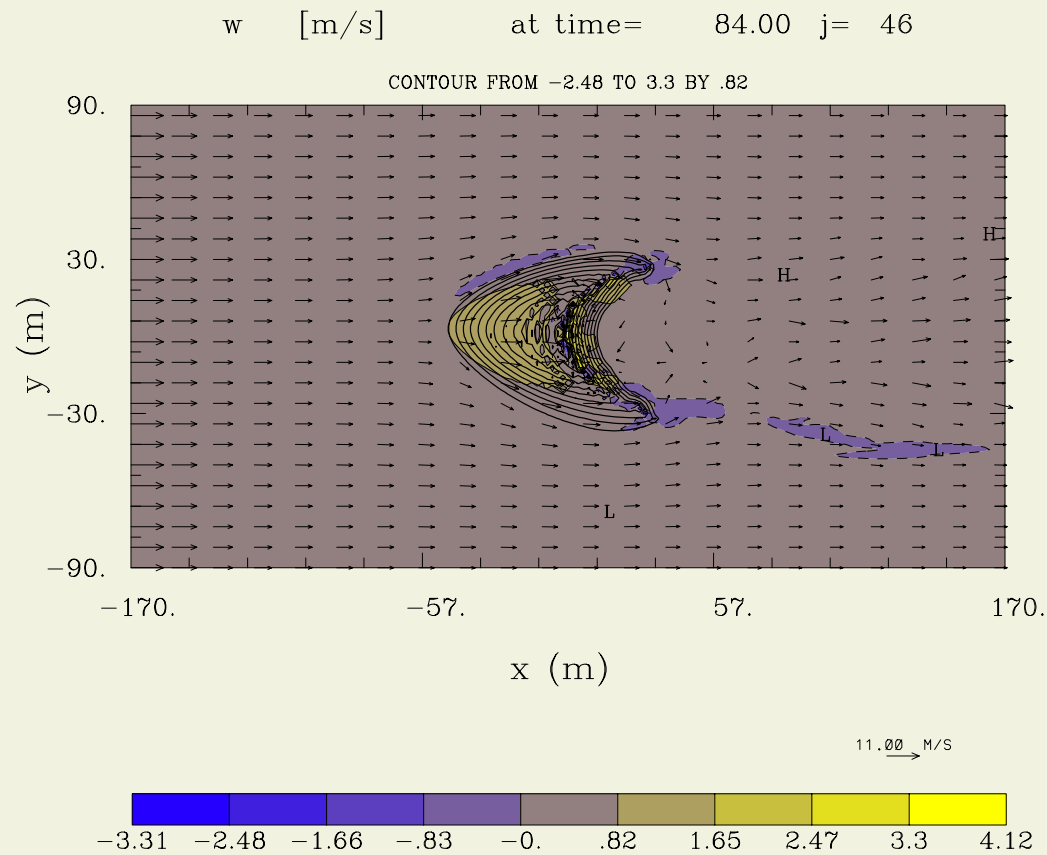
Non erodible substrate profile



Non erodible substrate xz

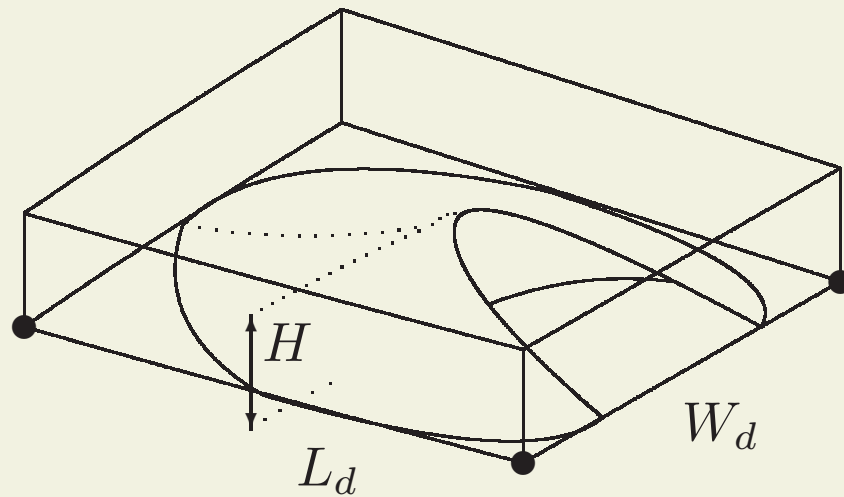


Non erodible substrate xy

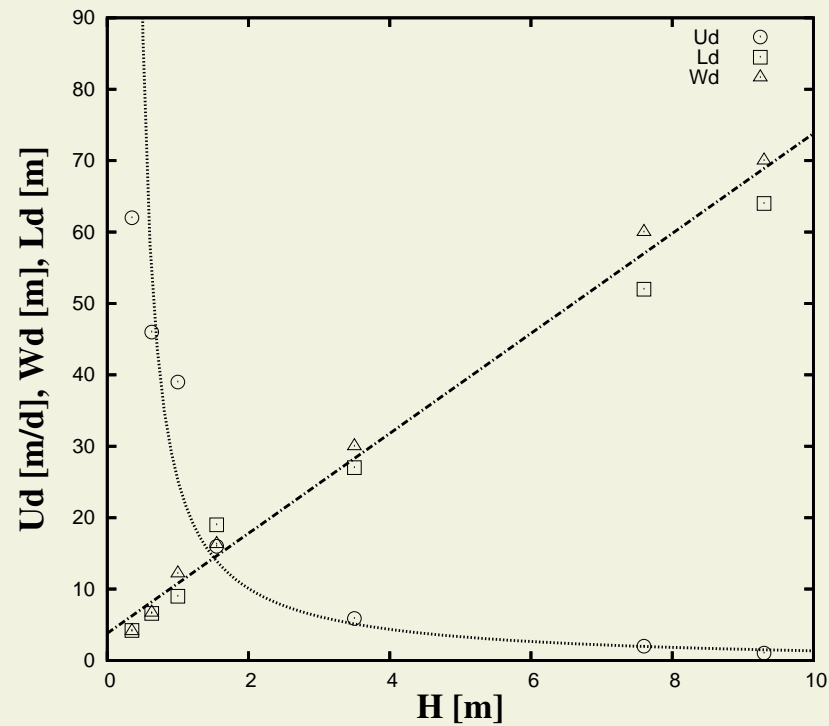


Sketch of a dune. Dimensions

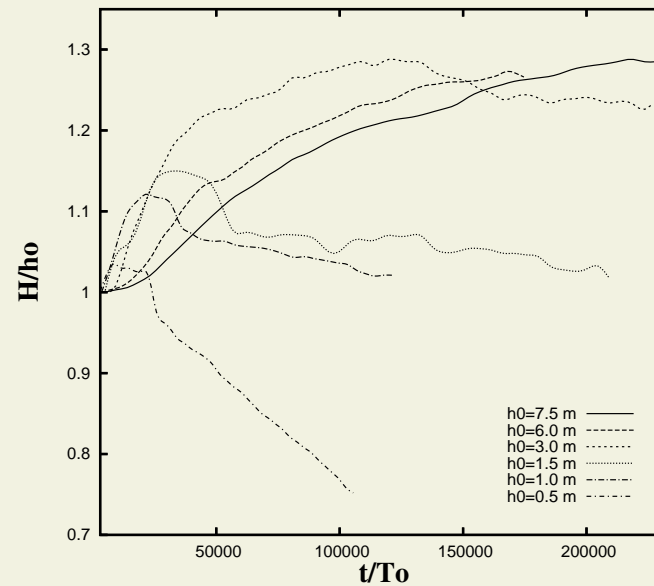
H : height, W_d : maximal width, L_d : maximal length



Non erodible substrate. Dune velocity, width and length



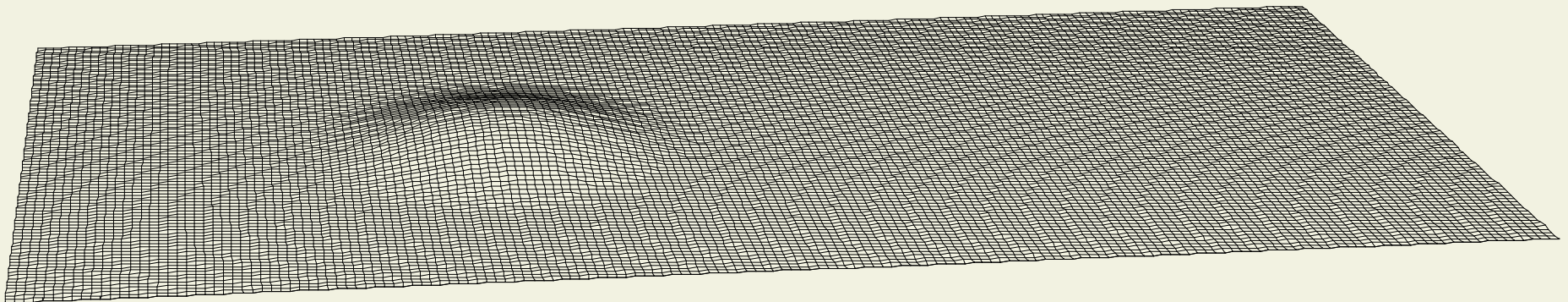
Maximum height



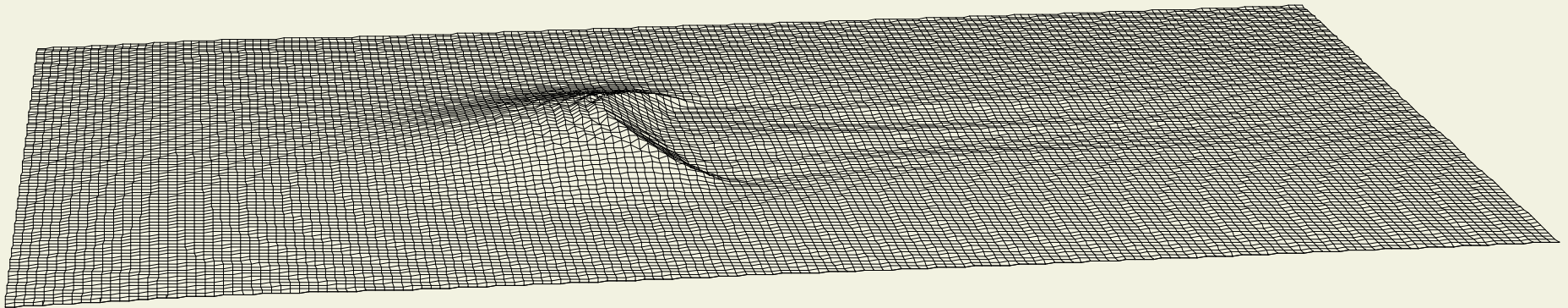
h_0 : Initial pile height, $T_0 = a / |\mathbf{v}_e|$.

- ★ Stabilization of final heights for $H > 1$ m and monotone decrease of heights for $H < 1$ m (match observations!)

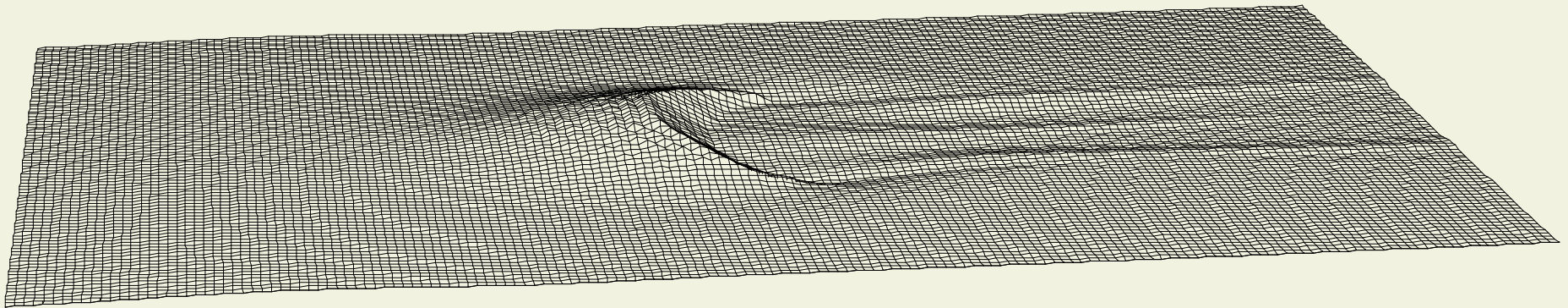
Erodible basement. Profile



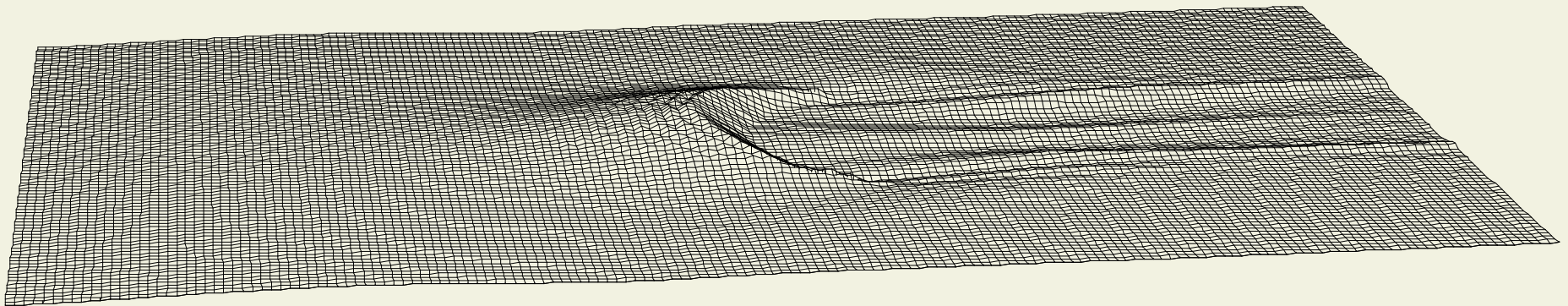
Erodible basement. Profile



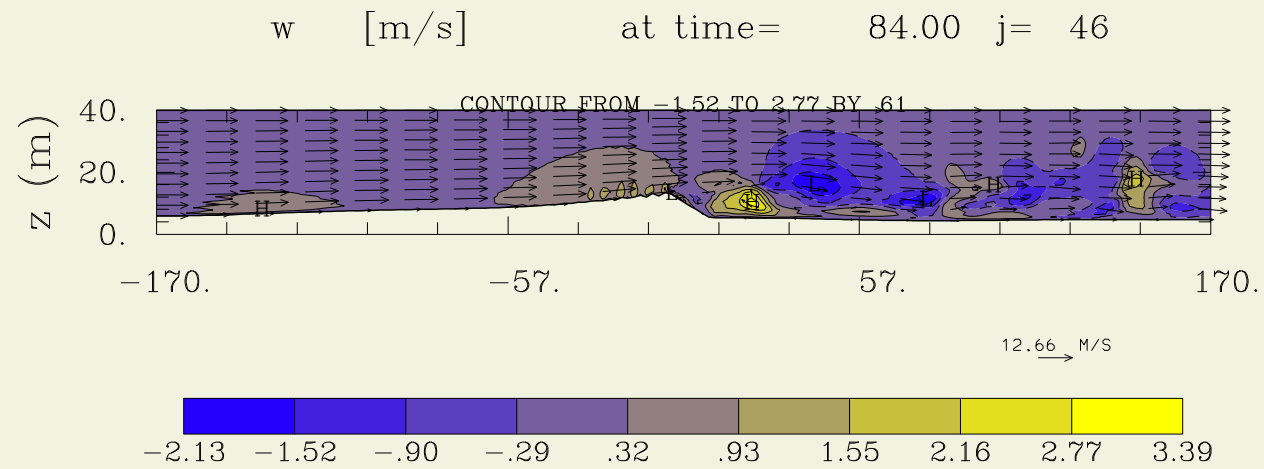
Erodible basement. Profile



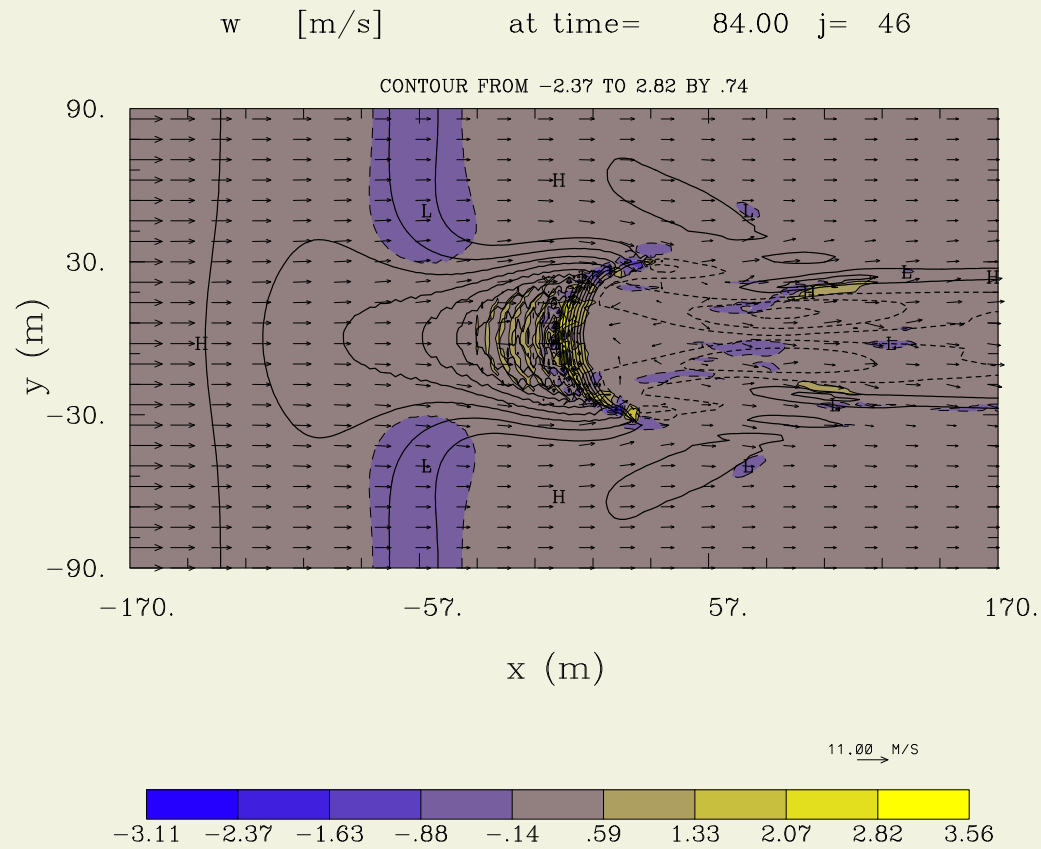
Erodible basement. Profile



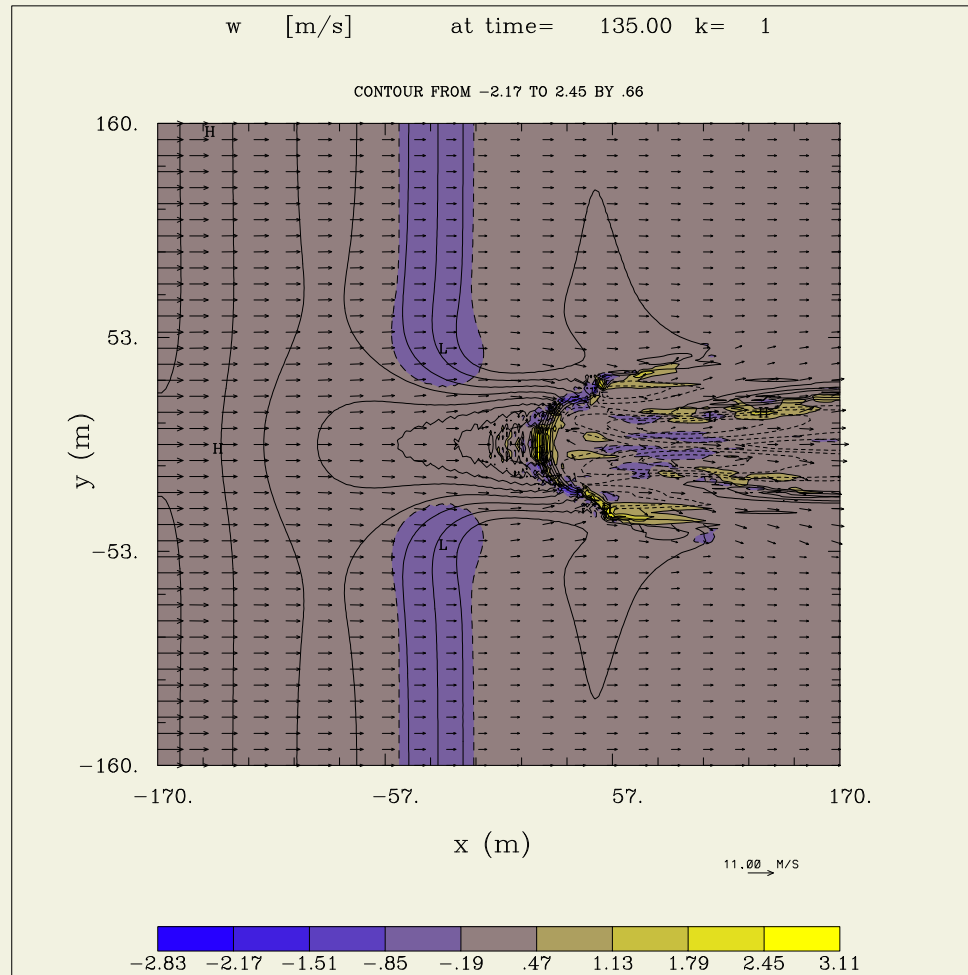
Erodible substrate xz



Erodible substrate xy



Erodible substrate xy: Long term

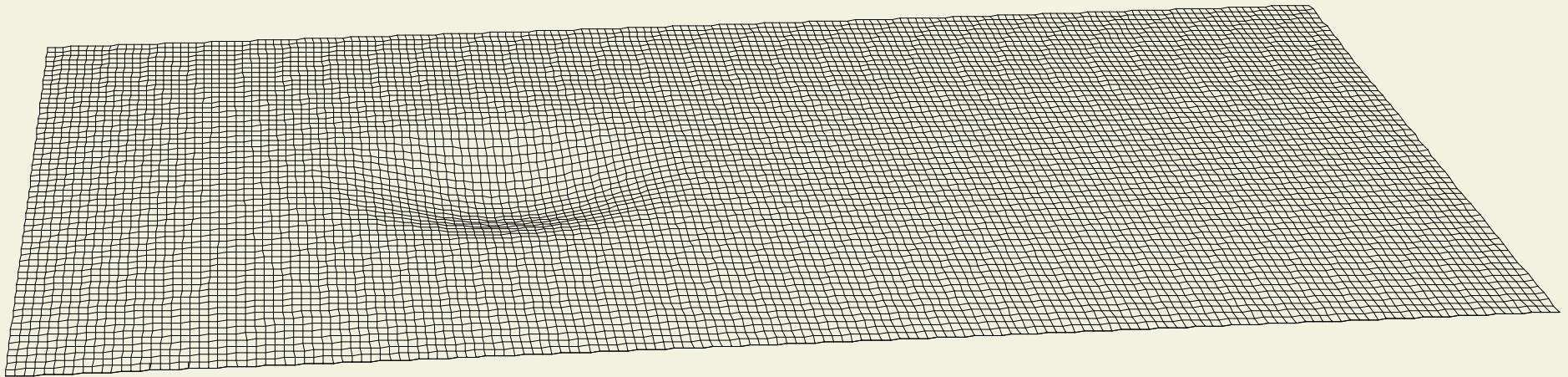


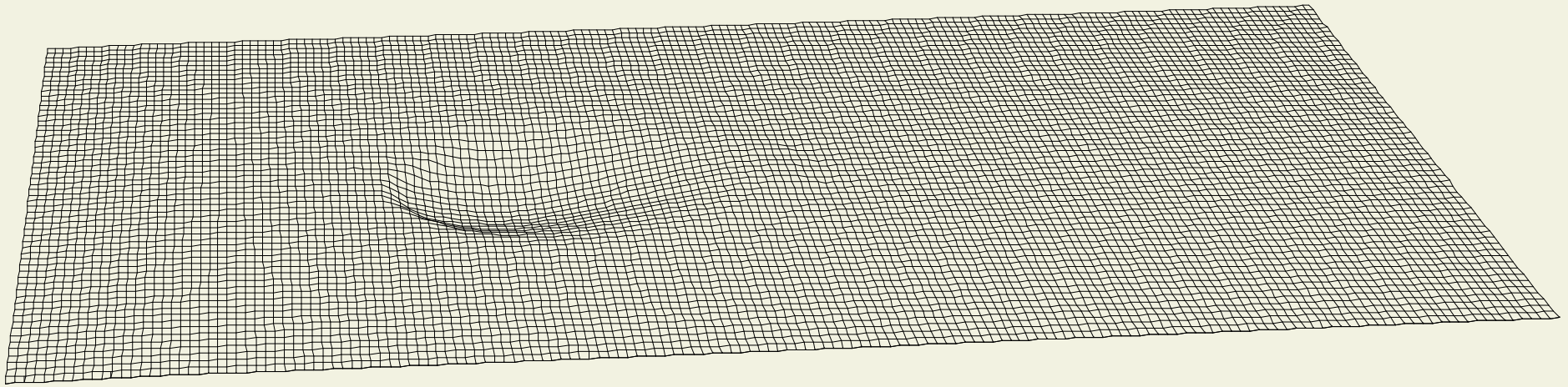
Remarks

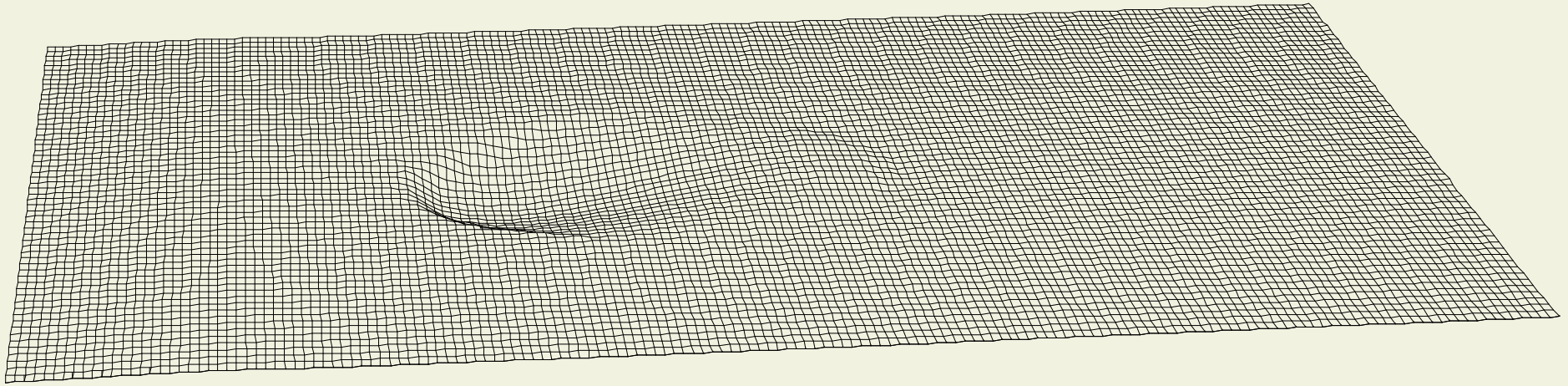
- Full coupled *severe wind scenario*: Numerical Challenge *per se*
Time dependent coordinate transformation, rescaling, LES, active
landform.

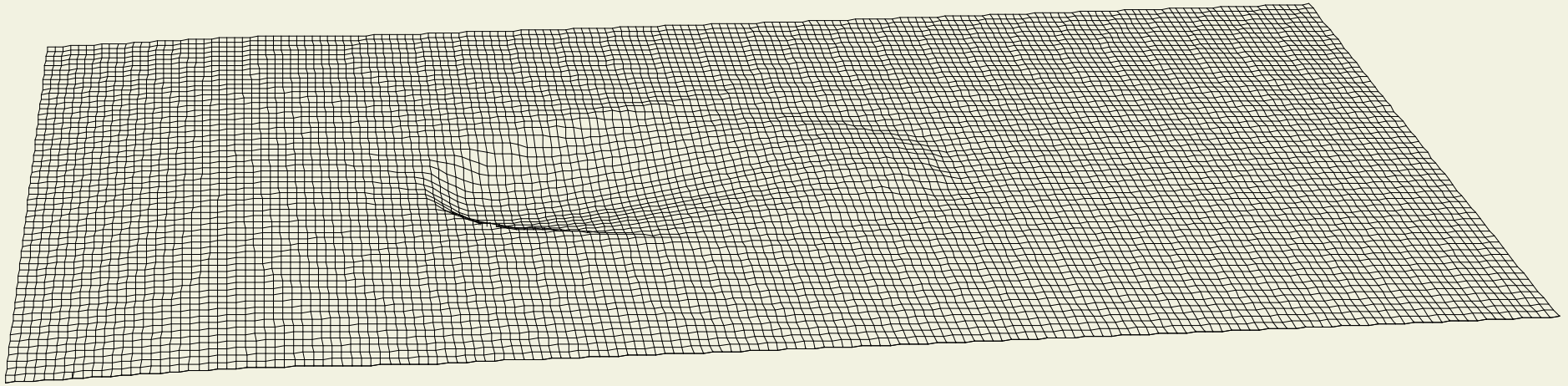
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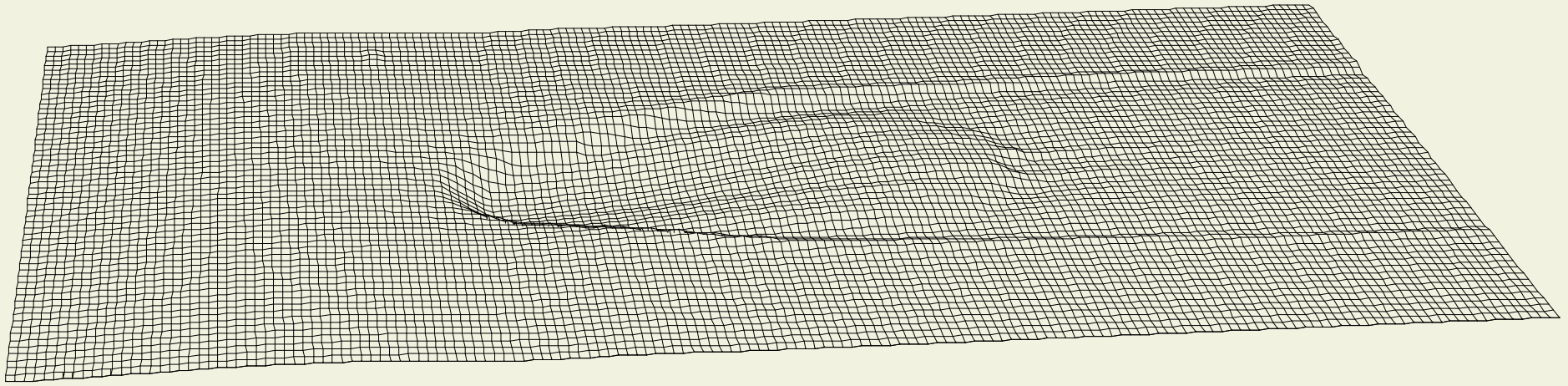
- Full coupled *severe wind scenario*: Numerical Challenge *per se*
Time dependent coordinate transformation, rescaling, LES, active landform.
- a) Dynamics of complex morphologies
 - Local scales: evolution of simple forms. Dependence on initial conditions

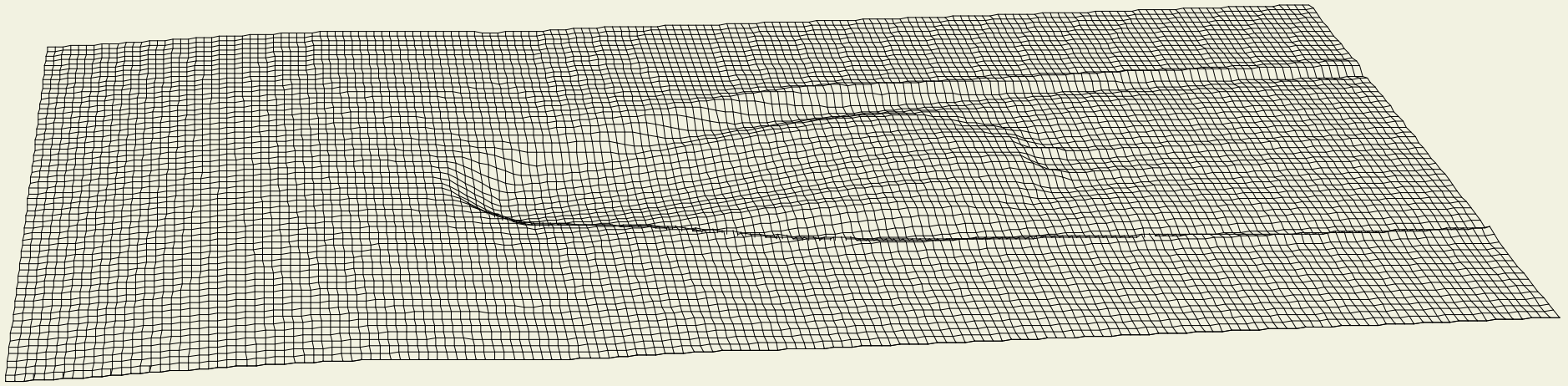


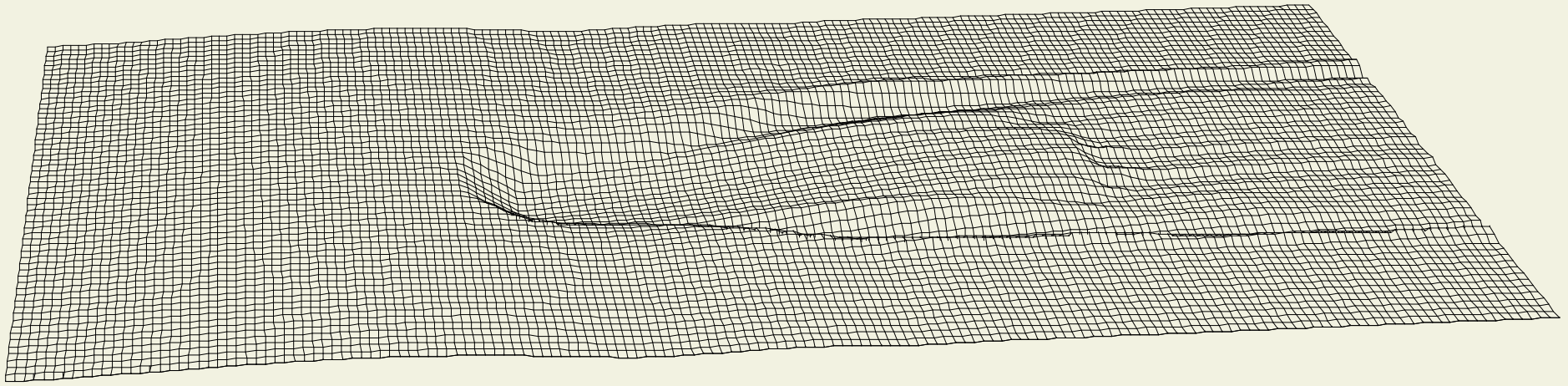








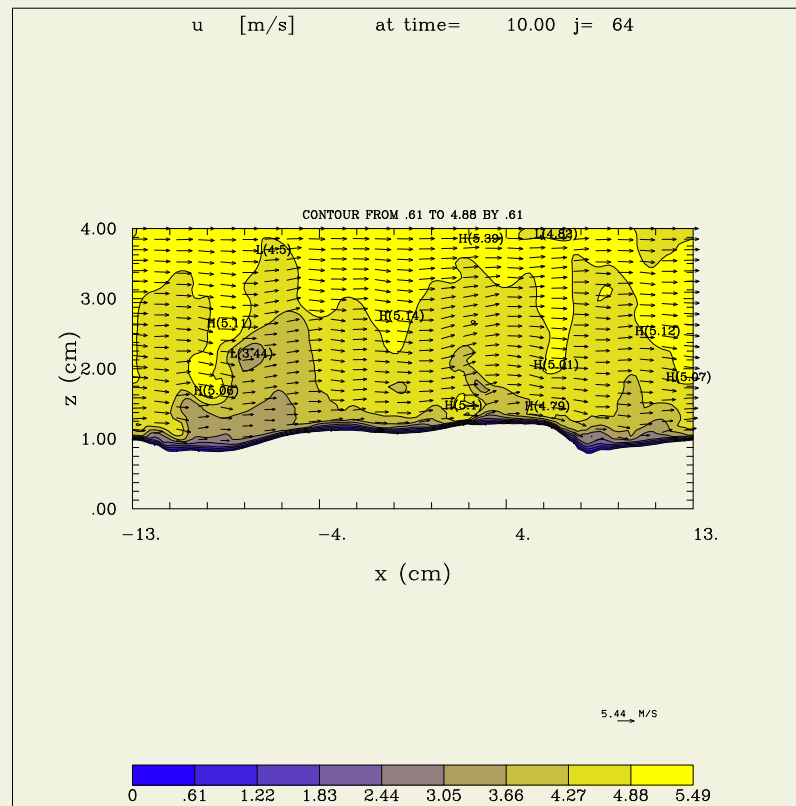




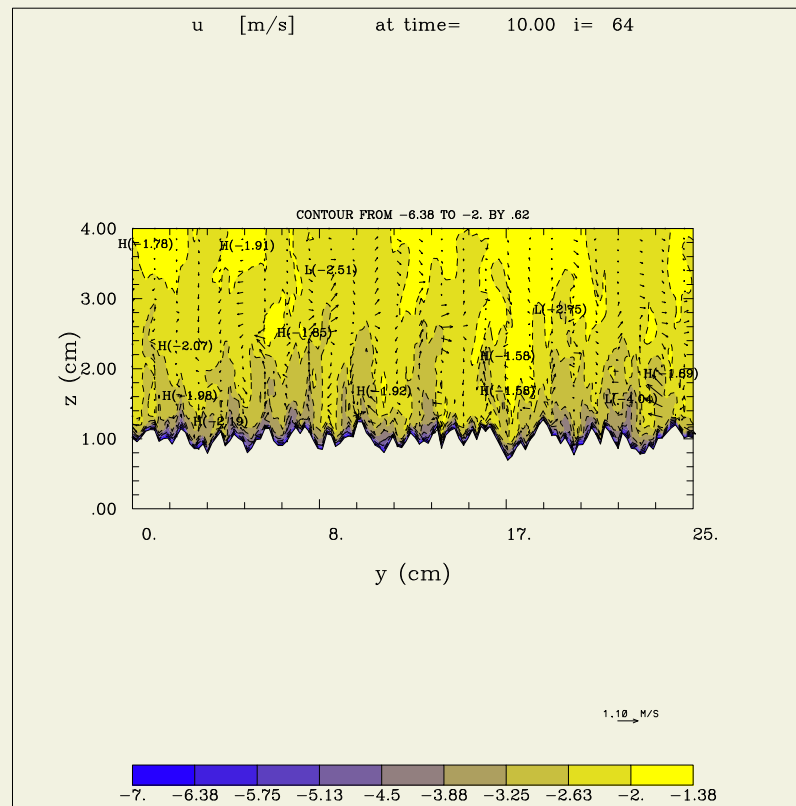
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 - Small scales: ripples

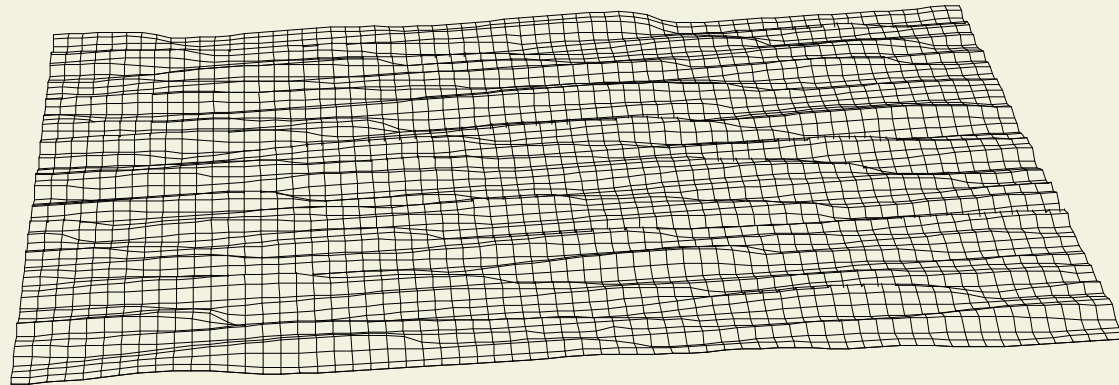
Ripple formation



Ripple formation: stripped structures



Ripple formation: stripped structures



Remarks

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 - Planetary scale

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- a) Dynamics of complex morphologies
 - Local scales: evolution of simple forms. Dependence on initial conditions
 - Small scales: ripples
 - Planetary scale
- b) Dynamics of more complex flows:
 - Rotating
 - Stratified and thermally forced
 - Free surface flows